Image-based modelling of ocean surface circulation from satellite acquisitions

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Objective

- Estimation of surface circulation (2D motion w(x, t)) from an image sequence T(x, t)
- Data Assimilation

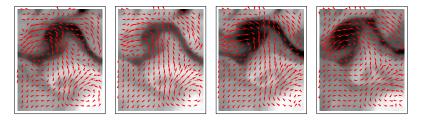


Figure : Satellite data acquired over the Black Sea. Circulation

• Important issue for pollutant transport and meteorology forecast



- Image structures as tracer of the sea surface circulation
- Optical flow methods:
 - They compute translational displacement between two observations
 - They are ill-posed (the Aperture Problem) and required spatial regularization
 - Validity of brightness constancy assumption?
 - \Rightarrow not physically suited for ocean circulation



- Circulation: advanced 3D oceanographic models are available (Navier-Stokes, see NEMO project (http://www.nemo-ocean.net) for instance)
- But: only a thin upper layer of ocean is observable (from satellite)
- From 3D Navier-stokes and various simplifications, a 2D ocean surface circulation model is derived → shallow water equations
- Compute an optimal solution w.r.t. the model and fitting observations: use of **Data Assimilation** techniques
- Need of an observation model: link between state vector and observations



- Shallow water model requires information on temperature and upper layer thickness
- Temperature: available from Sea Surface Temperature images (NOAA-AHVRR sensors)
- Layer thickness: not available from remote sensing
- $\Rightarrow\,$ We propose a method to compute surface circulation from SST image and without information on the upper layer thickness
- $\Rightarrow\,$ Use of a rough model, missing information will be represented in an additional model
- $\Rightarrow\,$ Solution is computed using a weak 4D-Var formulation



Model

Shallow water equations

State vector is velocity, **w** = (u, v)^T, surface temperature (T_s) and upper layer thickness (η):

$$\frac{\partial u}{\partial t} = -u \frac{\partial u}{\partial x} - v \frac{\partial u}{\partial y} + fv - g' \frac{\partial \eta}{\partial x} + K_{\mathbf{w}} \Delta u \qquad (1)$$

$$\frac{\partial \mathbf{v}}{\partial t} = -u\frac{\partial \mathbf{v}}{\partial x} - v\frac{\partial \mathbf{v}}{\partial y} - fu - g'\frac{\partial \eta}{\partial y} + K_{\mathbf{w}}\Delta v \qquad (2)$$

$$\frac{\partial \eta}{\partial t} = -\frac{\partial(u\eta)}{\partial x} - \frac{\partial(v\eta)}{\partial y} - \eta \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y}\right)$$
(3)
$$T_{s} = \partial T_{s} - \partial T_{s} + V \wedge T$$
(4)

$$\frac{s}{t} = -u\frac{\partial T_s}{\partial x} - v\frac{\partial T_s}{\partial y} + K_T \Delta T_s$$
(4)

- K_w , K_T are diffusive constants, f the Coriolis parameter and g' the reduced gravity (see paper for details)
- Functions w, T_s and η are defined on a space-time domain: $\Omega \times [0,T], \Omega \subset \mathbb{R}^2$

• Geophysical forces (in red) are grouped

$$\frac{\partial u}{\partial t} = -u\frac{\partial u}{\partial x} - v\frac{\partial u}{\partial y} + fv - g'\frac{\partial \eta}{\partial x} + K_{w}\Delta u$$
(5)
$$\frac{\partial v}{\partial t} = -u\frac{\partial v}{\partial x} - v\frac{\partial v}{\partial y} - fu - g'\frac{\partial \eta}{\partial y} + K_{w}\Delta v$$
(6)
$$\frac{\partial \eta}{\partial t} = -\frac{\partial(u\eta)}{\partial x} - \frac{\partial(v\eta)}{\partial y} - \eta\left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y}\right)$$
(7)
$$\frac{\partial T_{s}}{\partial t} = -u\frac{\partial T_{s}}{\partial x} - v\frac{\partial T_{s}}{\partial y} + K_{T}\Delta T_{s}$$
(8)

in a hidden part, $\mathbf{a} = (a_u, a_v)^T$, named "additional model"



• The previous system is rewritten as:

$$\frac{\partial u}{\partial t} = -u\frac{\partial u}{\partial x} - v\frac{\partial u}{\partial y} + a_u$$
(9)
$$\frac{\partial v}{\partial t} = -u\frac{\partial v}{\partial x} - v\frac{\partial v}{\partial y} + a_v$$
(10)
$$\frac{\partial T_s}{\partial t} = -u\frac{\partial T_s}{\partial x} - v\frac{\partial T_s}{\partial y} + K_T \Delta T_s$$
(11)

with

$$a_{u} = f_{v} - g' \frac{\partial \eta}{\partial x} + K_{w} \Delta u$$
(12)
$$a_{v} = -f_{u} - g' \frac{\partial \eta}{\partial y} + K_{w} \Delta v$$
(13)



where η verifies Eq.(3)

- The additional model **a** is now considered as an unknown that we want to retrieve
- The state vector is $\mathbf{X} = \begin{pmatrix} \mathbf{w} & T_s \end{pmatrix}^T$, and the model ruling the evolution in time of \mathbf{X} is summarized as:

$$\frac{\partial \mathbf{X}}{\partial t}(\mathbf{x},t) + \mathbb{M}(\mathbf{X})(\mathbf{x},t) = \begin{pmatrix} \mathbf{a}(\mathbf{x},t) \\ 0 \end{pmatrix} \quad \mathbf{x} \in \Omega, t \in (0,T]$$
(14)

with

$$\mathbb{M}(\mathbf{X}) = \begin{pmatrix} u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} \\ u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} \\ u \frac{\partial T_s}{\partial x} + v \frac{\partial T_s}{\partial y} + K_T \Delta T_s \end{pmatrix}$$



Initialization and Observation model

• Need of an initial condition for X(0):

$$\mathbf{X}(\mathbf{x},0) = \mathbf{X}_b(\mathbf{x}) + \varepsilon_B(\mathbf{x}), \quad \mathbf{x} \in \Omega$$
 (15)

 ε_B Gaussian with covariance matrix B

- Observations are SST images, $T(t_i)$, available at some given dates t_1, \cdots, t_N
- The observation operator $\mathbb H$ projects the state vector in the observation space. It is defined as:

$$\mathbb{H}(\mathbf{X}) = T_s \tag{16}$$

• Link between state vector and observation:

$$H(\mathbf{X})(\mathbf{x},t_i) = T(\mathbf{x},t_i) + \varepsilon_R(t_i), \quad \mathbf{x} \in \Omega, i = 1, \cdots, N$$
(17)

 ε_R Gaussian with covariance matrix R



Data assimilation Weak 4D-Var formulation

• To solve:

$$\frac{\partial \mathbf{X}}{\partial t}(\mathbf{x}, t) + \mathbb{M}(\mathbf{X})(\mathbf{x}, t) = \begin{pmatrix} \mathbf{a}(\mathbf{x}, t) \\ 0 \end{pmatrix} \mathbf{x} \in \Omega, t \in (0, \mathsf{T}]$$
(18)
$$\mathbb{H}(\mathbf{X})(\mathbf{x}, t_i) = T(\mathbf{x}, t_i) + \varepsilon_R(t_i), \quad \mathbf{x} \in \Omega, i = 1, \cdots, \mathsf{N}$$

$$\mathbf{X}(\mathbf{x}, 0) = \mathbf{X}_b(\mathbf{x}) + \varepsilon_B(\mathbf{x}), \quad \mathbf{x} \in \Omega$$

we minimize the cost function J:

$$J(\varepsilon_B, \mathbf{a}(t_1), \cdots, \mathbf{a}(t_N)) = \langle \varepsilon_B, B^{-1} \varepsilon_B \rangle + \gamma \sum_{i=1}^N \|\nabla \mathbf{a}(t_i)\|^2 + \sum_{i=1}^N \langle \mathbb{H}(\mathbf{X})(t_i) - \mathcal{T}(t_i), R^{-1}(\mathbb{H}(\mathbf{X})(t_i) - \mathcal{T}(t_i) \varepsilon_B \rangle$$
(19)

under the constraint of Eq.(18)

 $\bullet~\gamma$ term is introduced to prevent numerical instabilities



Weak 4D-Var formulation

Computation of ∇J

Theorem 1

Let $\lambda(\mathbf{x}, t)$ be an auxiliary variable (named adjoint variable) as solution of:

$$\lambda(T) = 0 \tag{20}$$

$$-\frac{\partial\lambda(t)}{\partial t} + \left(\frac{\partial\mathbb{M}}{\partial\mathbf{X}}\right)^* \lambda = \mathbb{H}^T R^{-1}[\mathbb{H}\mathbf{X}(t) - T(t)] \quad t = t_i \quad (21)$$
$$= 0 \quad t \neq t_i \quad (22)$$

Then, gradient of J is:

∂a

$$\frac{\partial J}{\partial \varepsilon_B} = 2 \left(B_I^{-1} \varepsilon_B + \lambda(0) \right)$$

$$\frac{\partial J}{\partial J} = 2 \left(a_I - b_I + \lambda(0) \right)$$
(23)

$$\frac{\partial S}{\langle t_i \rangle} = 2\left(-\gamma \Delta \mathbf{a}(t_i) + \lambda(t_i)\right) \tag{24}$$



Algorithm

- **9** Forward pass: integrate forward in time X(t), compute J
- **2** Backward pass: integrate backward in time $\lambda(t)$, compute ∇J
- Perform a steepest descent using numerical solver and get new values for X(0) and a(t_i), i = 1 ··· N
- Repeat steps 1, 2, 3 up to convergence



- Eq. (18) must be discretized:
 - In time: Euler scheme
 - In space:
 - components *u* and *v* are transported by a non linear advection (Burger equations): Godunov scheme
 - component I_s is transport by a linear advection: a first order up-wind
- Adjoint model: operator $\left(\frac{\partial \mathbb{M}}{\partial \mathbf{X}}\right)^*$ in Eq. (21) is formally defined as a dual operator:

$$\left\langle \phi, \left(\frac{\partial \mathbb{M}}{\partial \mathbf{X}}\right)^* \psi \right\rangle = \left\langle \left(\frac{\partial \mathbb{M}}{\partial \mathbf{X}}\right) \phi, \psi \right\rangle$$

It must be determined from the discrete model \mathbb{M}_{dis} : we use an automatic differentiation software, Tapenade [Hascoët and Pascual, 2013]

• Steepest descent is performed by BFGS solver [Byrd et al., 1995]



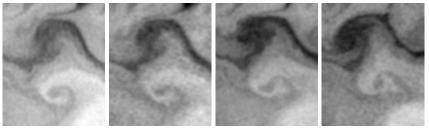
- Model: temperature diffusion is neglected ($K_T = 0$)
- Initial condition $\mathbf{X}_b = \begin{pmatrix} \mathbf{w}_b & I_b \end{pmatrix}^T$:
 - no information available on initial velocity, we set $\mathbf{w}_b = \vec{0}$
 - I_b is initialized to the first available observation $I(t_1)$
- Covariance matrix B: without information on initial velocity we choose $\varepsilon_B = \begin{pmatrix} 0 & 0 & \varepsilon_{B_{l_s}} \end{pmatrix}$ and we have:

$$\left\langle \varepsilon_{B}, B^{-1}\varepsilon_{B} \right\rangle = \left\langle \varepsilon_{B_{I_{s}}}, B_{I_{s}}^{-1}\varepsilon_{B_{I_{s}}} \right\rangle$$

- Matrix *B*_{*l*₅}: chosen diagonal, each element is set to 1 (1 Celsius degree means 25 % of image dynamics)
- Covariance matrix R: chosen diagonal, each element is set to 1
- γ : empirically fixed



 A sequence of four SST images was acquired over Black Sea on October 10th, 2007



(a) 30*min* (b) 6*h* (c) 15*h* (d) 30*h*

Figure : October 10th 2007, over Black Sea



Results Velocity retrieved

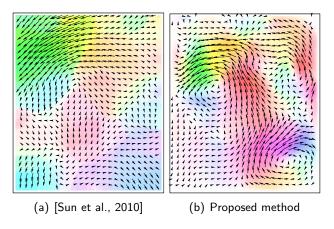


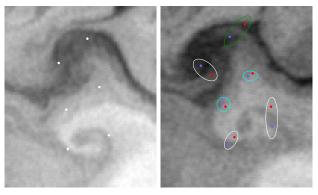
Figure : Motion computed between the first and second observation



- No ground-truth on satellite data, how to evaluate?
- We propose to compute the trajectory of some characteristic points in order to evaluate algorithms in term of transport
- A comparison with state-of-the-art is performed
- We proceed as follow:
 - Manual selection of a characteristic point in the first observation
 - A map of signed distance is computed
 - Distance map is transported by the velocity field that we want to evaluate [Lepoittevin et al., 2013]
 - Computation of local maximum in transported map gives the characteristic point position along the sequence



Result Characteristic points



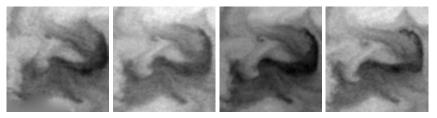
(a) First observation

(b) Last observation. Blue = our method, red = Sun et al.



Observations

• Sequence acquired on October 8th, 2005



(a) 30*min* (b) 10*h*30*min* (c) 12*h* (d) 15*h*30*min*

Figure : October 8th 2005, over Black Sea



Motion results

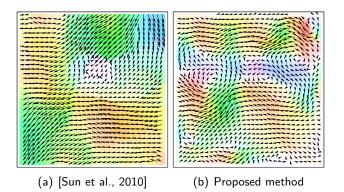
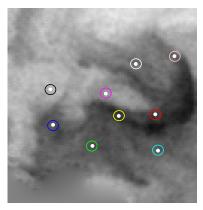


Figure : Motion computed between the first and second observation



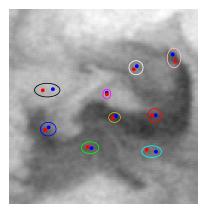
Characteristic points



(a) First observation



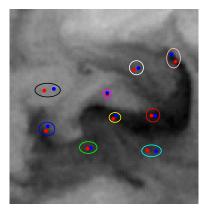
Characteristic points



(a) Second observation



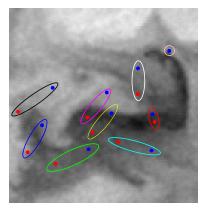
Characteristic points



(a) Third observation



Characteristic points

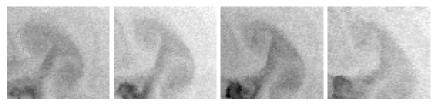


(a) Last observation



Observations

• Sequence acquired on July 27th, 2007



(a) 30*min*

(b) 8*h*15*min*

(c) 13*h*

(d) 22h30min



(e) 24*h*30*min*



Motion results

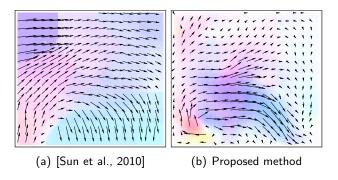
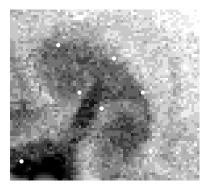


Figure : Motion computed between the first and second observation



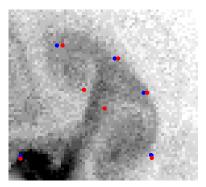
Characteristic points



(a) First observation



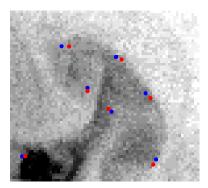
Characteristic points



(a) Second observation



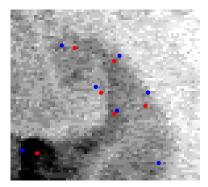
Characteristic points



(a) Third observation



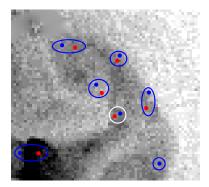
Characteristic points



(a) Fourth observation



Characteristic points

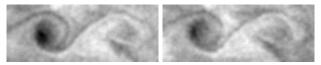


(a) Last observation



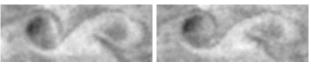
Observations

• Sequence acquired on May 14th, 2005



(a) 30*min*

(b) 2h55min



(c) 5*h*15*min*

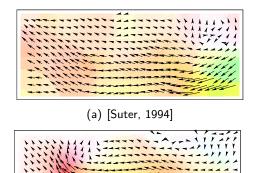
(d) 7h15min



(e) 16h15min



Motion results



(b) Proposed method

Tenner .

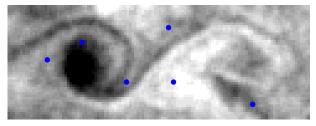
11117

-+++

Figure : Motion computed between the first and second observation



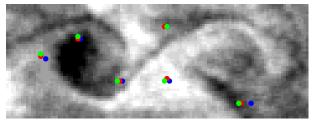
Characteristic points



(a) First observation



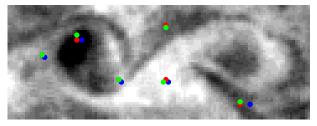
Characteristic points



(a) Second observation



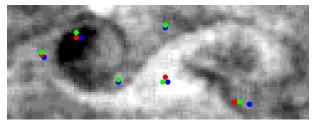
Characteristic points



(a) Third observation



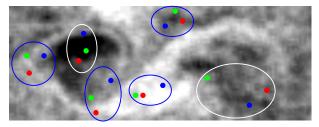
Characteristic points



(a) Fourth observation



Characteristic points



(a) Last observation



- \bullet Determination of surface circulation using a rough model ${\rm M}$
- \bullet Dynamics not modelled by ${\mathbb M}$ is captured in ${\boldsymbol a}$
- Analysis of a retrieved and comparison with a shallow water model
- Experiments on synthetic models (Temperature and upper layer thickness) with ground truth



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